國立中央大學地球科學系 GP3087 Sedimentary Geology

Instructor: 林殿順

This course introduces the principles of sedimentology and stratigraphy. Topics will be covered are: origin and transport of sedimentary materials, physical properties of sedimentary rocks, composition, classification, and diagenesis of sedimentary rocks, depositional environments, and principles of stratigraphy and basin analysis. This course includes a two-day field training in SW Taiwan.

Course Outline

Part I: Origin and Transport of Sedimentary Materials

- 1. Weathering and soils
- 2. Transport and deposition of siliciclastic sediments

Part II: Physical Properties of Sedimentary Rocks

- 3. Sedimentary textures
- 4. Sedimentary structures

Part III: Composition, Classification, and Diagenesis of Sedimentary Rocks

- 5. Siliciclastic sedimentary rocks
- 6. Carbonate sedimentary rocks
- 7. Other chemical/biochemical and carbonaceous sedimentary rocks

Part IV: Depositional Environments

- 8. Continental (terrestrial) environments
- 9. Marginal-marine environments
- 10. Siliciclastic marine environments
- 11. Carbonate and evaporite environments

Part V: Stratigraphy and Basin Analysis

- 12. Lithostratigraphy
- 13. Sequence and magnetic stratigraphy
- 14. Biostratigraphy
- 15. Chronostratigraphy and geologic time
- 16. Basin Analysis, Tectonics, and Sedimentation

Textbook

Boggs, S., Jr., 2006, *Principles of Sedimentology and Stratigraphy*, fourth edition: Prentice Hall, New Jersey, 662 pp.

Course Schedule

Class: 15 September ~ 3 November 2022; Midterm exam: 10 November 2022

Class: 17 November 2022 ~ 29 December 2022; Final exam: 5 January 2023

Field trip: 10 ~ 11 December 2022 (Southwest Taiwan, 曾文溪、七股潟湖)

Final exam: 05 January 2023

Grading: Midterm exam 35%; Final exam 35%;

Field trip 20%; Assignments 10%

Chapter 1. Weathering and Soils



Mechanical Chemical
How sediments are made and landscapes formed

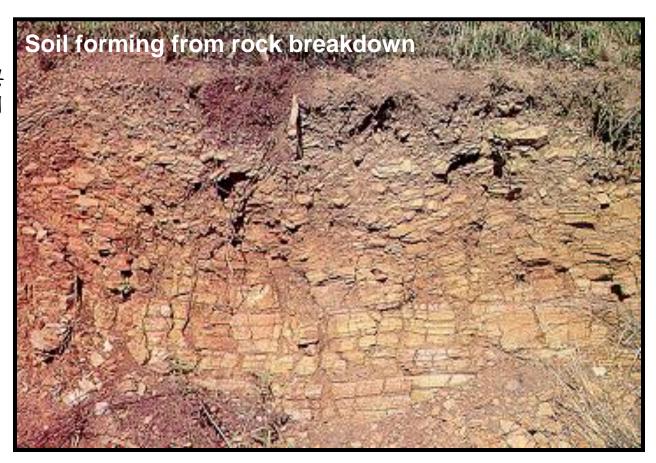
1.1 Introduction

Weathering

岩石因物理崩解、化學與生物作用而形成殘餘的固態物質與溶解(dissolved)的化學物質的過程稱為風化作用。

Regolith(表岩屑)

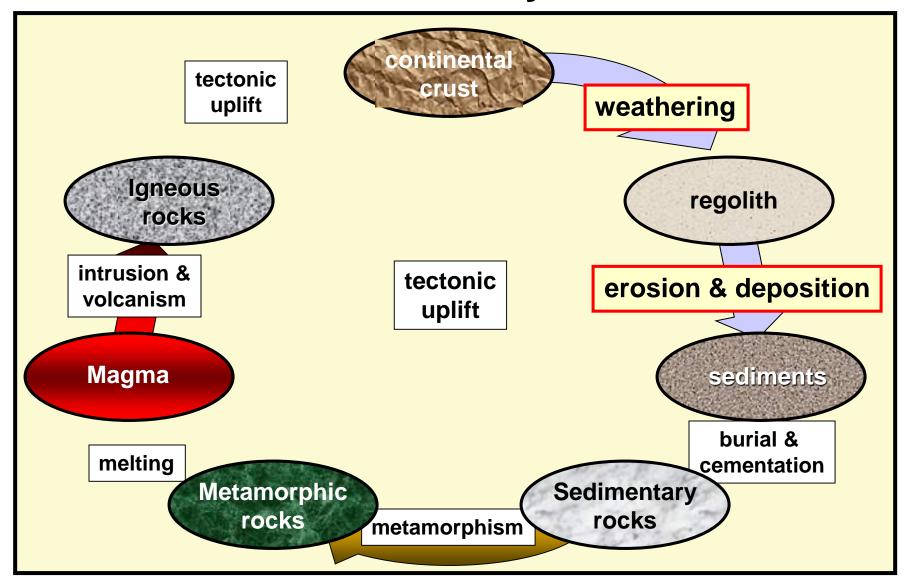
The loose layer of broken rock and mineral fragments on the surface.



Erosion

Transport of regolith and the wearing down of bedrock

The Rock Cycle

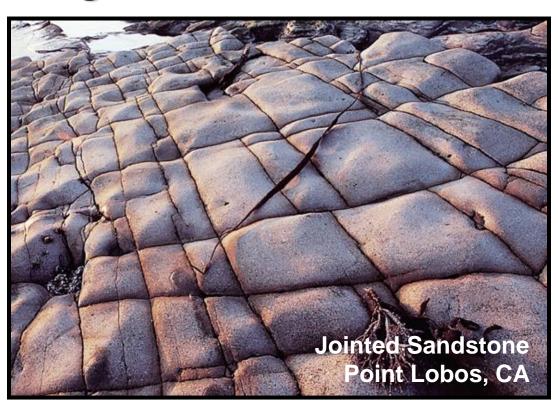


1.2 Subaerial weathering processes

1.2.1 Physical Weathering

The physical breaking of rocks into small particles

- Freeze-thaw (frost) weathering (ice wedging)
- Isolation (thermal) weathering
- Salt weathering
- Wetting and drying
- Stress-release weathering



Most mechanical weathering exploits JOINTS, pre-existing fractures in the rock

Freeze-thaw cycles: Ice wedging

The physical breaking of rocks into small particles

Water creeps into cracks during warm weather; it then expands as it freezes (about 9%)



This puts great pressure on the rocks

Ice wedging is important:

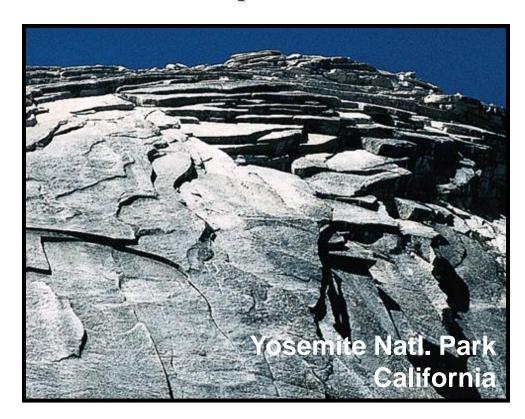
- Arctic climates
- Temperate climates
- High altitudes (mountains)

Fig. 1.1 Large, angular blocks of rock generated by **freeze-thaw weathering** of thin-bedded sandstones and mudstones of the Canning formation (Paleocene) exposed along the Canning River, Arctic National Wildlife Refuge, Alaska.

Release of overburden pressure

Large, thin sheets of rock that fall off bedrock

Upward rebound of the rocks creates tensile stresses, causing fractures to develop that are oriented nearly parallel to the topographic surface.





Spheroidal weathering is smaller-scale weathering of roughly massive rock masses, cut by intersecting joints, causing layers or "skins" to spall off to produce spheroidal cores.

Fig. 1.2 Spheroidal weathering in granite. Note low successive, thin layers of weathered rock are spalled off to produce a spheroidal core.

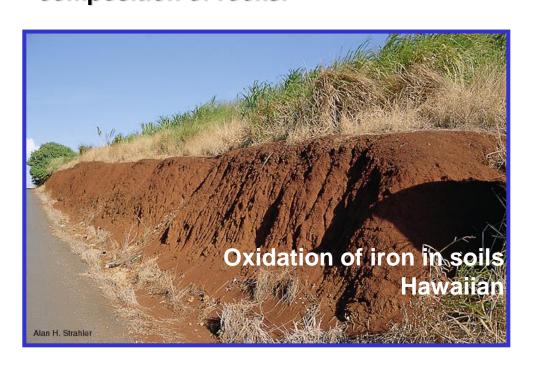


Weathering exploits joint surfaces: Example from Penghu basalts

1.2.2 Chemical Weathering

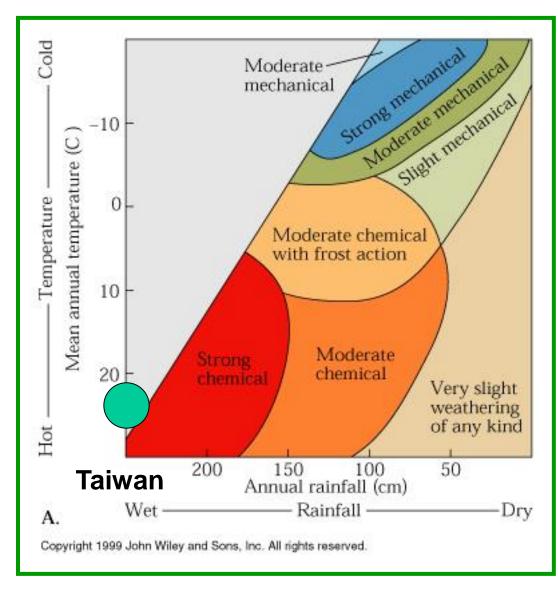
Chemical weathering alters both the chemical and mineralogical composition of rocks.





- A. Simple solution (congruent dissolution)
- B. Hydrolysis (水解)
- C. Oxidation (氧化)
- D. Others: Hydration (水合), Ion exchange (離子交換), Chelation

Chemical Weathering: Climate



Climate affects chemical weathering strongly

- More water
- More vegetation
- More soils
- Hotter Temperatures

Happens in any oxygenated environment (air, most water)

台灣平均雨量為250 cm/yr

Chemical weathering products vs. climate

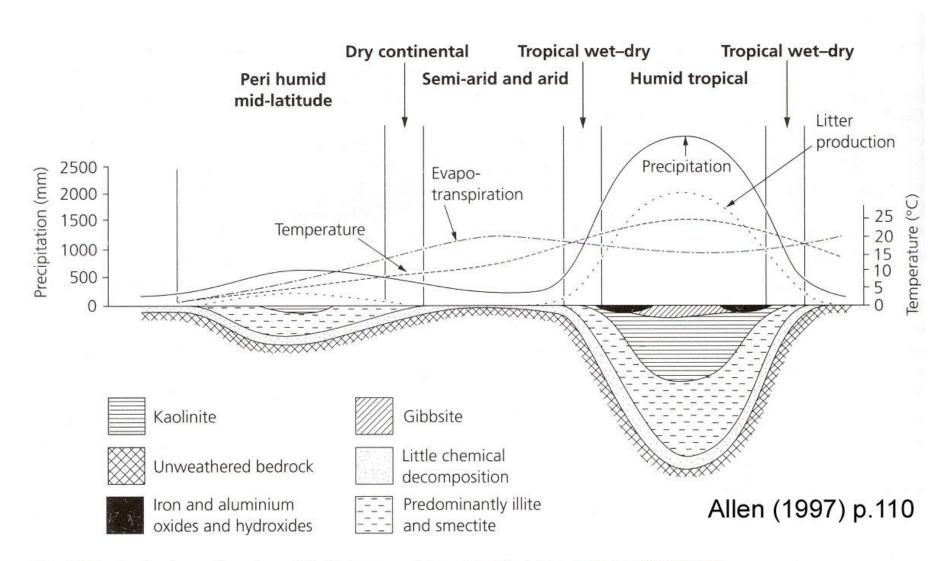
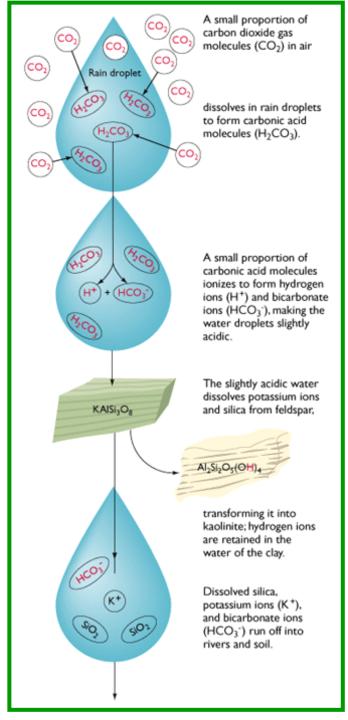


Fig. 3.9 Latitudinal zonation of regolith thickness and type. Modified from Strakhov (1967) [10].

Table. 1.1 Principal processes of chemical weathering

| Most important processes | Examples | Principal kinds of rock materials affected | |
|--|---|--|--|
| Simple (congruent) Solution—Dissolution of soluble minerals in H_2O (direct solution) or in $H_2O + CO_2$ (carbonation) to | $SiO_2 + 2H_2O \rightarrow H_4SiO_4$ (direct solution) (quartz) (silicic acid) aq | Highly soluble minerals (e.g., gypsum, halite), quartz | |
| yield cations and anions in solution | CaCO ₃ + H ₂ O + CO ₂ \leftrightarrow Ca ²⁺ + 2HCO ₃ ⁻ (Carbonation) (calcite) aq aq | Carbonate rocks | |
| Hydrolysis (incongruent dissolution)— Reaction between H ⁺ and OH ⁻ ions of water and the ions of silicate minerals, | $2KAlSi_3O_8 + 2H^+ + 9H_2O \rightarrow H_4Al_2Si_2O_9 + 4H_4SiO_4 + 2K^+$ (orthoclase) aq (kaolinite) (silicic acid) aq | Silicate minerals | |
| yielding soluble cations, silicic acid, and clay minerals (if Al present) | 2NaAlSi ₃ O ₈ + 2H ⁺ + 9H ₂ O \rightarrow H ₄ Al ₂ Si ₂ O ₉ + 4H ₄ SiO ₄ + 2Na ⁺ (albite) aq (kaolinite) (silicic acid) aq | | |
| Oxidation—Loss of an electron from an element (commonly Fe or Mn) in a mineral, resulting in the formation of oxides or | $2\text{FeS}_2 + 15/2\text{O}_2 + 4\text{H}_2\text{O} \rightarrow \text{Fe}_2\text{O}_3 + 4\text{SO}_4^{\ 2^-} + 8\text{H}^+$ (pyrite) (hematite) aq aq | Iron- and manganese-bearing silicate minerals, iron sulfides | |
| hydroxides (if water present) | $MnSiO_3 + 1/2O_2 + 2H_2O \rightarrow MnO_2 + H_4SiO_4$ (rhodonite) (pyrolusite) (silicic acid) | | |
| Other Processes | | | |
| Hydration and Dehydration—Gain (hydration) or loss (dehydration) of water molecules from a mineral, | $Fe_2O_3 + H_2O \leftrightarrow 2FeOOH \text{ (hydration)}$ (hematite) (goethite) | Ferric oxides | |
| resulting in formation of a new mineral | $CaSO_4 \cdot 2H_2O \leftrightarrow CaSO_4 + 2H_2O$ (dehydration) (gypsum) (anhydrite) | Evaporites | |
| Ion Exchange —Exchange of ions, principally cations, between solutions and minerals | K -clay + Mg^2 + \leftrightarrow Mg -clay + K^+ Ca-zeolite + Na^+ \leftrightarrow Na -zeolite + Ca^{2+} | Clay minerals and zeolites | |
| Chelation—Bonding of metal ions to organic molecules having ring structures | Metal ions (cations) + chelating agent (e.g., secreted by lichens) \rightarrow H ⁺ ions + chelate (metal ions/organic molecules in solution) | Silicate minerals | |



B. Hydrolisis

水解作用指由氫(H+)或氫氧根(OH-)離子取代原來礦物的其他元素(如K+, Na+, Ca²+, Mg²+),而形成新的礦物。水解作用有CO₂加入,則稱為碳酸化作用(Carbonation)。

NEEDS WATER TO OCCUR!

 CO_2 和水作用產生H+與 HCO_3 -形成弱酸性

2KAISi₃O₈ (orthoclase)+2H⁺+9H₂O \longrightarrow Al₂Si₂ O₅(OH)₄ (kaolinite)+4H₄SiO₄(aq) +2K⁺

Weak acids strip cations out of minerals

This is the main process of:

- weathering silicates
- making clays

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Figure from Dr. Alan Jay Kaufman at http://www.geol.umd.edu/%7Ekaufman/kaufman.html

石灰岩的化學風化作用主要是碳酸化作用(Carbonation)

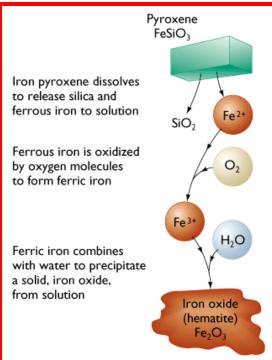




 $H_2CO_3 + 2CaCO_3 \longrightarrow Ca^{2+} + 2HCO_3^{-}$

若CO₂多則反應向右,石灰岩溶解, 形成卡斯特(karst)地形與鐘乳石洞或 滲穴(sinkhole);若CO₂少反應向左, 碳酸鈣沉澱,形成鐘乳石。在乾燥 氣候下,碳酸鈣於土壤中沉澱,形 成鈣質結礫岩(calcrete)。

C. Oxidation



Loss of an electron from an element (commonly Fe or Mn) in a mineral, resulting in the formation of oxides or, if water is present, hydroxides.

$$4FeSiO_{3} + O_{2} \longrightarrow 2Fe_{2}O_{3} + 4SiO_{2}$$
pyroxene hematite
$$Fe^{2+} \longrightarrow Fe^{3+} + e^{-}$$



D. Other chemical weathering processes: Hydration

某些礦物可將水分子(H₂O)納入其結晶格內,如此結晶內的孔隙加大,使此水化合物更容易受風化。另外,礦物體積膨脹之後,岩石的內部壓力加大,可以導致岩石破碎。



 $2Fe_2O_3 + 3H_2O \longrightarrow 2Fe_2O_3.3H_2O$ Hematite Limonite 赤鐵礦 褐鐵礦

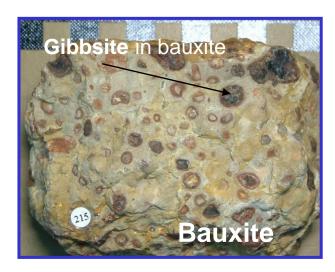
Products of oxidation

氧化物(Oxides)與氫氧化物(hydroxides)合稱為三氧化鋁鐵(sesquioxides)

| Si的氧化物 | 石英、非晶質矽石(amorphous silica)、蛋白石質矽石 (opaline silica) | |
|-----------------|---|----------|
| Al的氫氧化物 | Gibbsite(三水鋁石AI(OH) ₃ , boehmite(軟水鋁石γAI(OH) ₃) 一水硬鋁石(diaspore, αAIO(OH)) | Hematite |
| Fe的氧化物與氫 氧化物 | Hematite (赤鐵礦Fe ₂ O ₃), Magnetite (磁鐵礦,Fe ₃ O ₄), goethite(針鐵礦αFe(OH) | |
| Ti的氧化物 | Anatase (銳鈦礦,TiO ₂) | Goethite |

Bauxite (鋁土礦), the principal ore of aluminum, is a rock mixture consisting mostly of several hydrous aluminum oxide minerals, including boehmite, diaspore, gibbsite, and impurities such as QUARTZ, CLAY MINERALS, and iron hydroxides.

Example five Al-minerals that weather to form bauxite: kaolinite, muscovite, biotite, hornblende, feldspar.



Chemical Weathering: Clay Minerals

Generated strictly as a weather product by many types of chemical weathering

KAISi₃O₈ + 2H⁺ + 12H₂O
$$\longleftrightarrow$$
 KAI₃Si₃O₁₀(OH) ₂ + 2K⁺ + 6H₄SiO₄
K-spar Muscovite Clay
Hydrolysis + Hydration

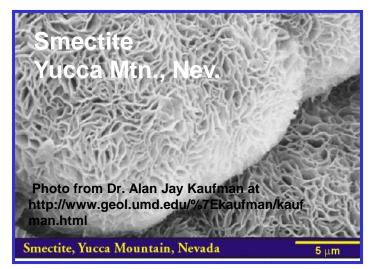
$$2KAI_3Si_3O_{10}(OH)_2 + 2H^+ + 3H_2O \longleftrightarrow 3AI_2Si_2O_5(OH)_4 + 2K^+$$
Muscovite Clay

Kaolinite Clay

Hydrolysis + Hydration

$$Al_2Si_2O_5(OH)_4 + 5H_2O \longleftrightarrow 2Al(OH)_3 + 2H_4SiO_4$$
Kaolinite Clay Gibbsite Clay

Hydrolysis + Hydration



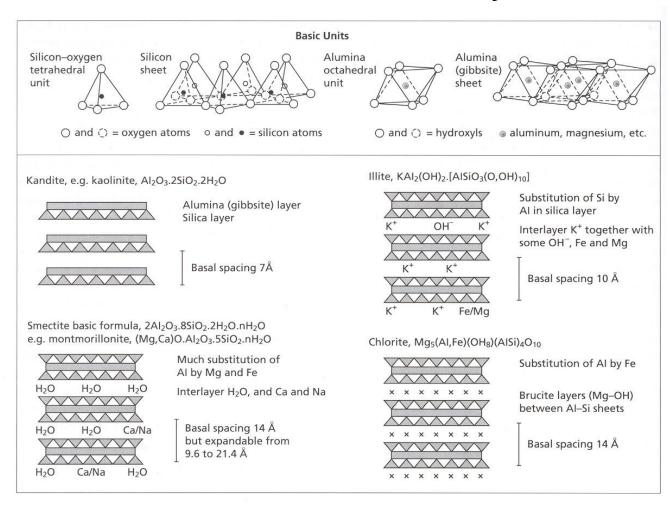


Clay Minerals

黏土礦物基本上是含氫氧根(OH-)或水 (H_2O) 之鋁矽酸鹽,其中有些礦物會有Mg, Fe取代AI的現象。屬葉狀矽酸鹽(phyllosilicates)。

| 礦物群 | 化學成分 | 主要礦物 | 四面體 (tetrahedral) 層:八面體 (octahedral) 層 | 附註 |
|-------------------|---|---|--|---|
| Kandite group | 含水、鋁矽 酸鹽 hydrous aluminium silicate | Kaolinite(高嶺石 $AI_2Si_2O_5(OH)_4$), halloysite (禾樂石 $AI_2Si_2O_5(OH)_4$.n H_2O) | 1:1 | |
| Illite group | 含水、鉀矽 酸鹽 | Illite(伊萊石(K,H ₃ O)(Al, Mg,Fe) ₂ (Si,Al) ₄ O ₁₀ (OH) ₂ .H ₂ O) | 2:1 | illite最常見的黏土礦 物 |
| Smectite group | 複合、含水、 鎂鋁矽酸鹽 | Montmorillonite(蒙脫石,大約化學式(AI,Fe, Mg, Zn, Ca, Na) $_4$ (Si $_4$ O $_{10}$) $_2$ (OH) $_4$.nH $_2$ O);vermiculite(蛭石,Mg $_3$ (Si, AI) $_4$ (OH) $_2$.4.5H $_2$ O[Mg] $_{0.35}$) | 2:1 | Expandable clay |
| Chlorite group | Hydrous silicate of AI, Fe and Mg | Chlorite (綠泥石,(Mg,Fe) ₃ (Si,Al) ₄ O ₁₀).(Mg,Fe) ₃ (OH) ₆) | 2:2 | 主要是非風化產物、 常見於變質岩(如綠色 片岩)與火成岩(Mg-Fe 矽酸鹽的換質產物) |
| Mixed-layer group | 混成黏土 | Illite-montmorillonite mixed-layers | | 20 |

Structures of clay minerals



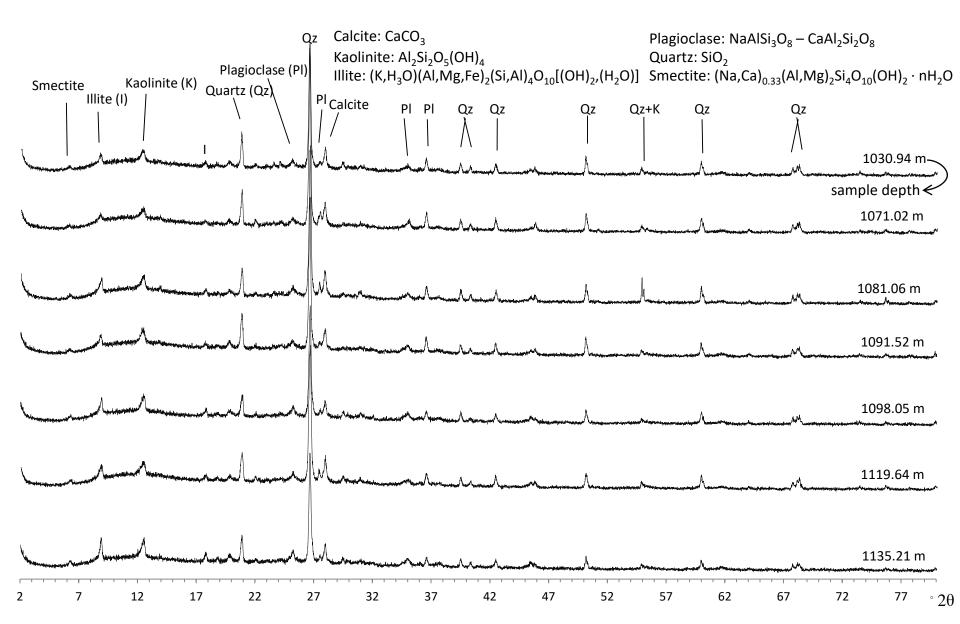




Tucker (2001) p.98

note: gibbsite layer: 由AI-O/OH組成的八面體層片稱之,因gibbsite(水鋁石) 的化學是為AI(OH)₃ Brucite layer: 由Mg-O/OH組成的八面體層片稱之,因brucite(水鎂石)化學式為Mg(OH)₂

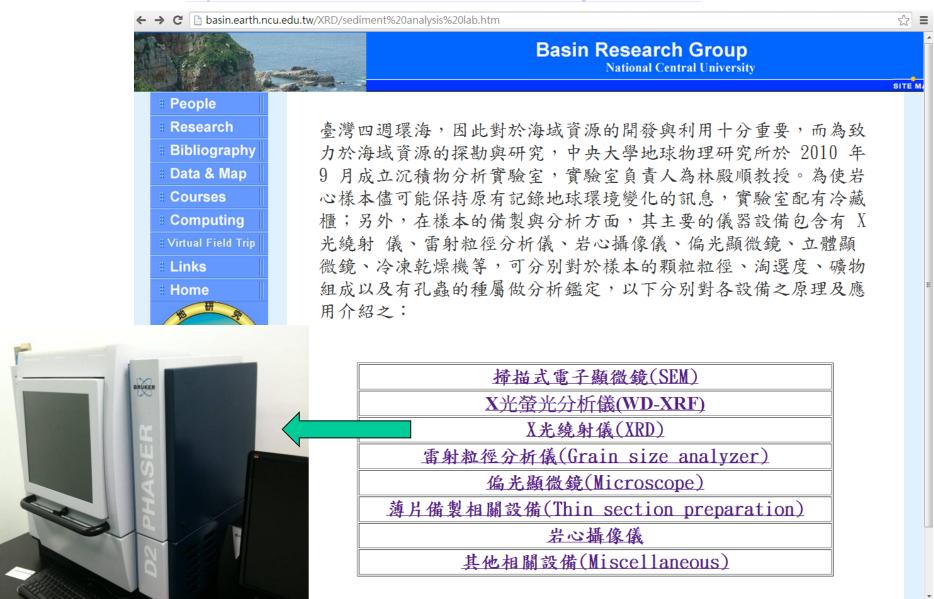
- 上圖Chlorite部分,x成一行表示brucite layer故為2:2
- 上圖層間的陽離子,很容易於風化或成岩過程中被leached out of the clay mineral by percolating waters.



XRD diffractograms of shale samples (bulk rock Chinshui Shale) from TCDP-A borehole, Taichung

國立中央大學地球科學系盆地研究室設備

http://basin.earth.ncu.edu.tw/XRD/sediment%20analysis%20lab.htm



Effects of Chemical Weathering on Basalt and Granite

Solid weathering Soluble ions Neutral species Minerals products in solution in solution Na¹⁺ and Ca²⁺ Feldspar — Clay minerals Silica **Basalt** Fe²⁺ mineral — Clay minerals and goethite Mg²⁺ Silica Magnetite — Goethite Na1+ and K1+ Feldspar → Clay minerals Silica Mica → Clay minerals Granite Fe²⁺ mineral → Clay minerals and goethite Mg²⁺ Silica

1.2.3 Weathering Rate

Goldich's weathering series vs. Bowen's reduction series:

Table 1.2 Relative stability of common sand-size minerals and various clay-size minerals under conditions of weathering

| Sand- and silt-size minerals* | | Clay-size minerals** | | |
|---|---|---|--|--|
| Sand- and sil Mafic minerals Olivine Pyroxene Amphibole Biotite | It-size minerals* Felsic minerals Ca plagioclase Ca-Na plagioclase Na-Ca plagioclase Na plagioclase K-feldspar, | 1. Gypsum, halite 2. Calcite, dolomite, apatite 3. Olivine, amphiboles, pyroxenes 4. Biotite 5. Na plagioclase, Ca plagioclase, K-feldspar, volcanic glass 6. Quartz 7. Muscovite 8. Vermiculite (clay mineral) 9. Smectite (clay mineral) | | |
| (Increasing stabilit | muscovite, quartz | 10. Pedogenic (soil) chlorite 11. Allophane (clay mineral) 12. Kaolinite, halloysite (clay minerals) 13. Gibbsite, boehmite (clay minerals) 14. Hematite, goethite, magnetite 15. Anatase, titanite, rutile, ilmenite (all titanium-bearing minerals), zircon | | |

unstable

stable

Source: *Goldich (1938), **Jackson (1968).

1.2.4 Products of subaerial weathering

Table. 1.3 Principal kinds of products formed by subaerial weathering processes and the types of sedimentary rocks ultimately formed from these products.

| Weathering process | Type of weathering product | Example | Ultimate depositional product | | |
|---------------------|----------------------------|---|--|--|--|
| Physical weathering | Particulate residues | Silicate minerals such as quartz and feldspar; all types of rock fragments | Sandstones, conglomerates mudrocks | | |
| Chemical weathering | | | | | |
| Hydrolysis | Soluble constituents | Silicic acid (H ₄ SiO ₄); K ⁺ , Na ⁺ , Mg ²⁺ , Ca ²⁺ , etc. | Cherts, limestones, etc. | | |
| | Secondary minerals | Clay minerals | Mudrocks (shales) | | |
| Simple solution | Soluble constituents | Silicic acid; K ⁺ , Na ⁺ , Mg ²⁺ , Ca ²⁺ , HCO ₃ ⁻ , SO ₄ ²⁻ , etc. | Limestones, evaporites, chert, etc. | | |
| Oxidation | Secondary minerals | Ferric oxides (Fe ₂ OOH); manganese oxides (MnO ₂) | Minor constituent in siliciclastic rocks | | |
| | Soluble constituents | Silicic acid; SO ₄ ²⁻ | Chert, evaporites, etc. | | |

Products of weathering: A. siliciclastic sedimentary rocks

矽質沉積岩占75%的地球 表面沉積物。

風化作用造成的固態物質,經常以石英為主,稱為矽質顆粒(siliciclastic grains),由這些顆粒造成的沉積岩稱為矽質沉積岩。矽質沉積岩包括礫岩、砂岩、頁岩等。

Note: 這些風化而來的固態物質也經常來自陸地, 所以他們也稱為陸源顆粒 (terrigenous grains)。

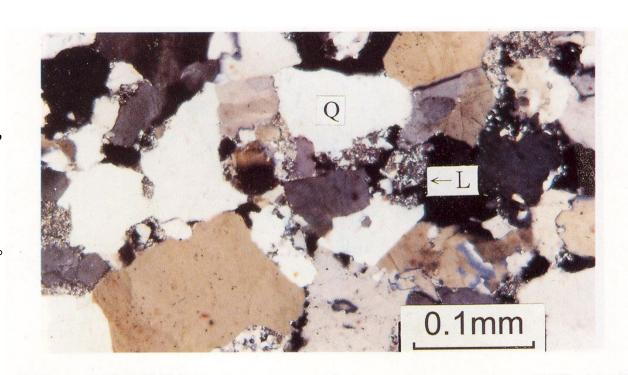


圖3-26 石英砂岩岩象

顯微鏡下的白冷層石英砂岩,由石英(Q)和少許岩屑(L)組成。顆粒多次角形,淘選度佳。由於受過深埋,所以顆粒間有些微的壓溶現象,形成齒狀穿插的邊緣(南投縣水里鄉)。

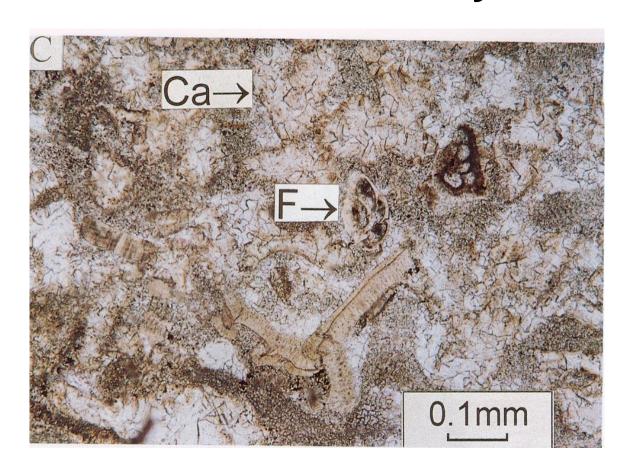
鄧屬予(1997)



Products of weathering: B. Chemical/biochemical sedimentary rocks

化學/生物化學沉積岩 占25%的地球表面沉積物。

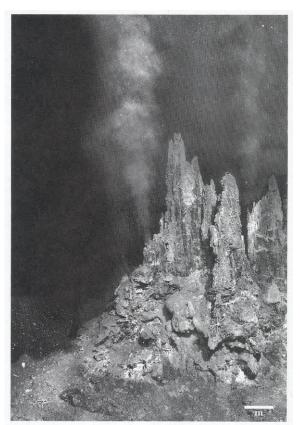
風化作用造成的溶解物(如Ca, Mg, Si)到達湖或海時,因化學或生物化學作用,將這些溶解物轉化成固態,形成沉積物。如石灰岩(limestone)、燧石(chert)、蒸發岩(evaporite)等。



Limestone (半屏山石灰岩)

鄧屬予(1997)

1.3 Submarine weathering processes and products



- 1. Interaction of seawater with hot oceanic rocks along mid-ocean ridge; Magnesium, sulfate, and sodium ions are removed from seawater, whereas elements such as iron, manganese, silicon, potassium, lithium, and strontium are enriched in the seawater. Products at oceanic hot springs: sulfide (pyrite, FeS2, chalcopyrite, CuFeS2), sulfate and oxides.
- Low-temperature (<20°C) alteration of volcanic rocks and sediments on the ocean floor. Products: smectite clay mineral, zeolite, chlorite.

Fig. 1.3 A multiple-orifice black smoker, Faulty Towers complex, Mothra hydrothermal vent field, Endeavour Segment, Juan de Fuca ridge. The constructional chimneys in the foreground were built by precipitation of sulfides and other minerals from heated water issuing from the vents at temperatures exceeding 250°C.

Submarine weathering process is important for changes in composition of seawater owing to ion exchange during the reaction of seawater with basalts.

1.4 Soils

1.4.1 Soil-forming processes

1. Additions to the ground surface - precipitation of dissolved ions in rainwater; influx of solid particles such as windblown dust; addition of organic matter from surface vegetation.

2. Transformations

- a. Decomposition of organic matter within soils to produce organic compounds.
- b. Weathering of primary minerals; formation of secondary minerals, including iron oxides.

3. Transfers

- Movement of solids or suspended material downward from one soil horizon to a lower horizon by groundwater percolation (eluviation).
- b. Accumulation of soluble or suspended material in a lower horizon (illuviation).
- Transfer of ions upward by capillary movement of water and precipitation of ions in the soil profile.
- 4. Removals: removal of substances still in solution to become part of the dissolved constituents in groundwater or surface water.
- 5. **Bioturbation of soil** soil disruption by animals and plants.

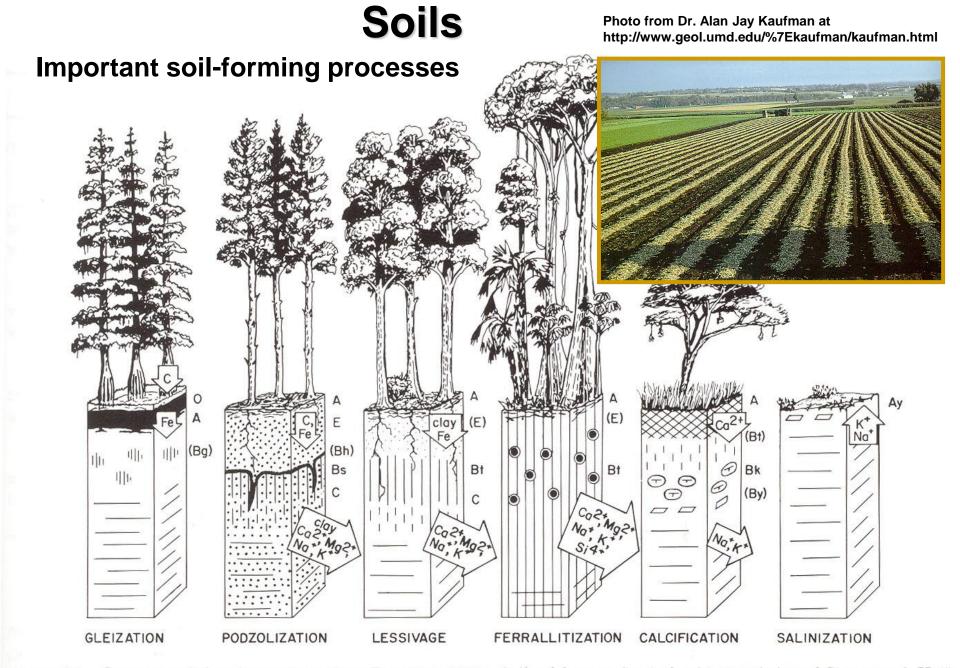
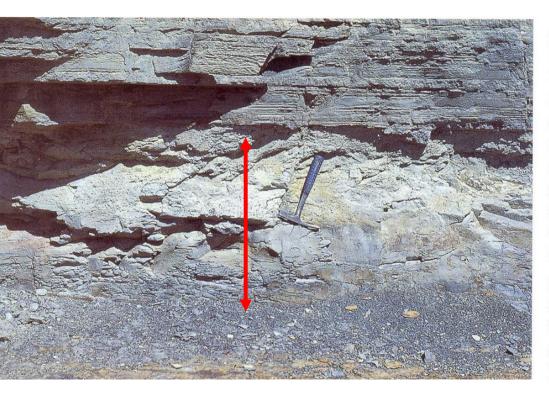


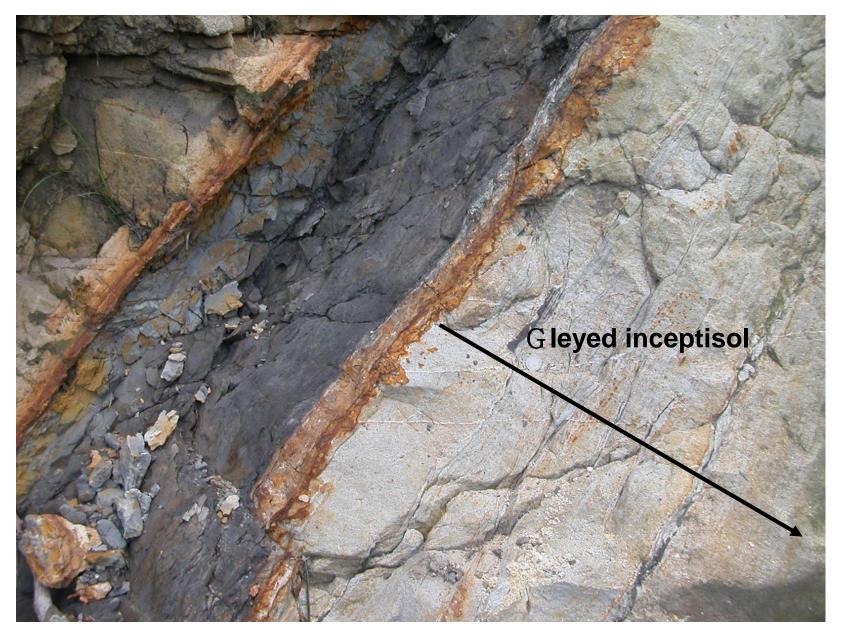
Figure 1.17. Common soil-forming regimes (from Retallack, 1990a, Soils of the past, fig. 4.12, with permission of Chapman & Hall)

Gleization (潛水灰化作用)

在缺(低)氧的泡水土壤中(waterlogged soil) ,使Fe³⁺還原成Fe²⁺ ,形成藍灰色或綠灰色土壤的作用。



Colour photo 100. Gleyed Inceptisol. The massive claystone below the bedded siltstones is a paleosol with abundant siderite and the blue-grey to purple-grey colour that is characteristic of waterlogged soils. Such colours and minerals of reducing chemical conditions are called gley features. Stagnant water in the profile has allowed preservation of much organic matter, including recognizable root traces and faecal pellets. However, it was not always waterlogged, as can be seen from deeply penetrating root traces, burrows like those of earthworms (see Colour photo 19), and profile differentiation into a light-coloured surface over a purple and clayey subsurface horizon. This horizon differentiation has not completely obscured relict bedding, so the paleosol was an Inceptisol, and in view of its gleyed nature, an Aquept. This is the Turimetta clay eroded phase paleosol in the Early Triassic, Garie Formation, south of Bilgola Beach, New South Wales, Australia (Retallack, 1976, 1977). Hammer handle for scale is 25 cm long.



東坑層(苗栗明德水庫)



東坑層(苗栗明德水庫)

Podzolization (灰化作用)

酸性淋溶作用(leaching),使表層的AI、Fe、有機質往下傳輸。本作用造成表層殘留石英,底層則聚集三氧化鋁鐵(sesquioxides)或有機質。

Lessivage (澱積作用)

將黏土由表層(A層)往下搬運至B層的裂隙澱積。本作用造成B層含較多的黏土,且顏色較表層為淺。

Calcification (鈣化作用)

在半乾燥到亞潮濕(subhumic)的氣候區內,碳酸鈣形成結核、聚集成帶,造成calcrete。

Ferrallitization (or laterization 紅鋁鐵礦化)

強烈與深部風化,造成厚層、成分均匀的土壤剖 面。土壤富含黏土與三氧化鋁鐵。



Near the front entrance of National Central University

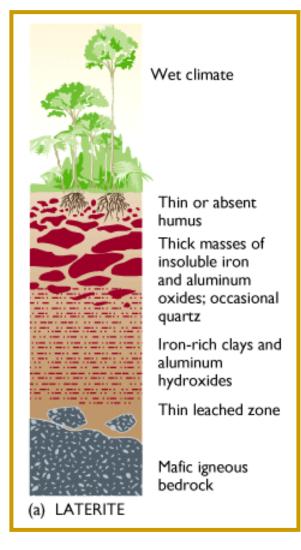
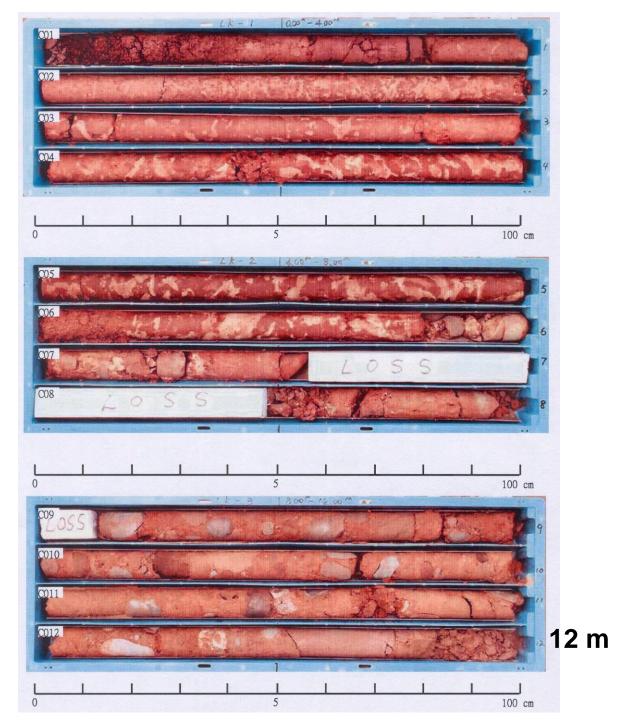


Figure from Dr. Alan Jay Kaufman at http://www.geol.umd.edu/%7Ekaufman/kaufman.html

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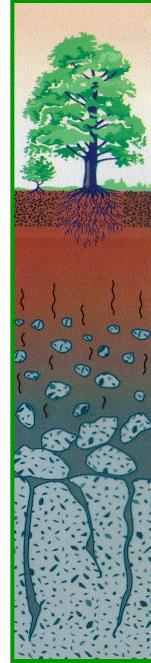
1.4.2 Soil profiles and soil classification

林口台地紅土剖面



Soil horizons





Soil Profile

0 horizon: 有機質層(e.g. coal)

A horizon: 表土層(屬洗出層) B horizon: 底土層(屬洗入層)

C horizon: 被岩層

Topsoil

Organic processes

Soil leached of soluble minerals; rich in clay and insoluble minerals and organic matter

Hydrolisis and Oxidation

Little organic mater; dissolved minerals from A horizon precipitated

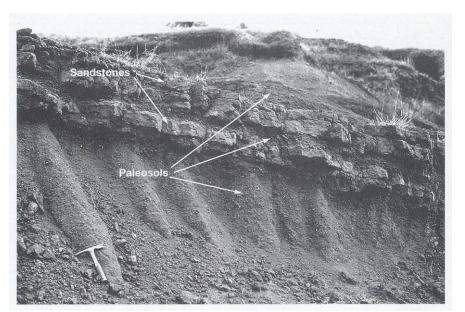
Hydrolisis New Minerals

Bedrock cracked and weathered

Mechanical weathering

39

1.4.3 Paleosols







Paleosols in the Miocene Penghu basalts

Fig. 1.7 Red paleosols exposed below bedded sandstones in the Middle Miocene, Chinji Formation, Siwalik Group, in a creek bed 3 km south of Khaur, Potwar Plateau, Pakistan. The hammer is 25cm long.

1.4.4 Recognition of paleosols

DIAGNOSTIC FEATURES OF PALEOSOLS

ROOT TRACES



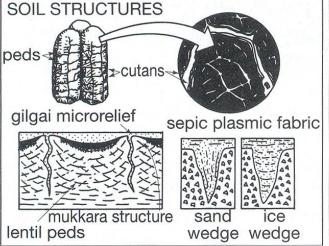
truncated tops
downward taper
downward branching

SOIL HORIZONS

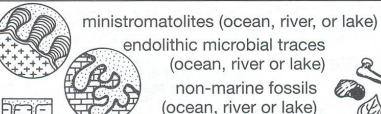


rip-up clasts in overlying sediment erosional, sharp top gradational changes downward

little altered parent material



PROMINENT OR COMMON FEATURES OF PALEOSOLS (and where else formed)



burrows (ocean, river or lake)



coal and carbonaceous shale (ocean, river or lake)
zones of base depletion (hydrothermal system)
quartz-rich residuum (ocean, river or lake)
zones of clay accumulation (ocean, river, lake,
deep burial, or hydrothermal system)
zones of carbonate accumulation (ocean, river, lake,
shallow or deep burial, or hydrothermal system)
zones of iron accumulation (ocean, river, lake,
or hydrothermal system)



nodules and concretions (ocean, river, lake, shallow burial, volcanic ash or hydrothermal system)



"desert roses" and crystals (playa lake, sabkha)



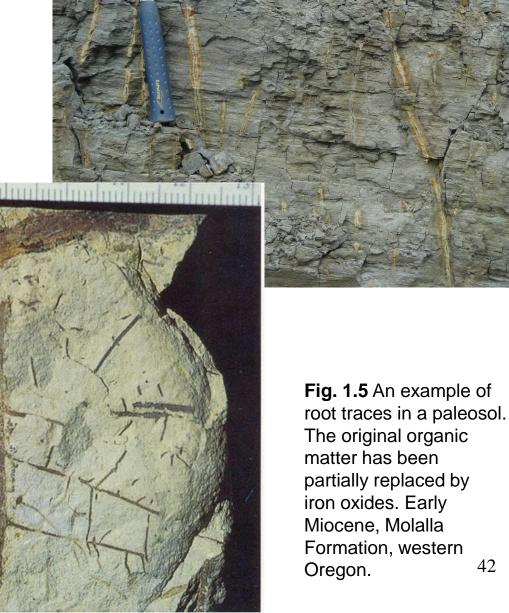
relict bedding (ocean, river, or lake)



relict crystal structure (playa lake, sabkha, shallow or deep burial, metamorphic, hydrothermal or igneous) relict foliation (fault zone, or metamorphic)

Fig. 1.4
Characteristic and common features useful in recognition of paleosols.

Root traces – one of the indicators for paleosol



卓蘭層(大安溪)

| TYPE | PLATY | PRISMATIC | COLUMNAR | ANGULAR BLOCKY | SUBANGULAR BLOCKY | GRANULAR | CRUMB |
|--------------------------|--|---|--|--|---|--|--|
| SKETCH | | | | | | *0.00 20.00 20.00 | ::: |
| DESCRIPTION | tabular and horizontal to land surface | elongate with flat top and vertical to land surface | elongate with domed top and vertical to surface | equant with sharp interlocking edges | equant with dull interlocking edges | spheroidal with slightly interlocking edges | rounded and spheroidal but not interlocking |
| SUAL HORIZON | E,Bs,K,C | Bt | Bn | Bt | Bt | A | Α |
| MAIN LIKELY CAUSES | initial disruption of relict bedding; accretion of cementing material | swelling and shrinking on wetting and drying | as for prismatic, but with greater erosion by percolating water, and greater swelling of clay | cracking around roots and burrows, swelling and shrinking on wetting and drying | as for angular blocky, but with more erosion and deposition of material in cracks | active bioturbation and coating of soil with films of clay, sesquioxides, and organic matter | as for granular; including fecal pellets and relict soil clasts |
| SIZE CLASS | very thin < 1 mm | very fine < 1 cm | very fine < 1 cm | very fine < 0.5 cm | very fine < 0.5 cm | very fine < 1 mm | very fine < 1 mm |
| | thin 1 to 2 mm | fine 1 to 2 cm | fine 1 to 2 cm | fine 0.5 to 1 cm | fine 0.5 to 1 cm | fine 1 to 2 mm | fine 1 to 2 mm |
| | medium 2 to 5 mm | medium 2 to 5 cm | medium 2 to 5 cm | medium 1 to 2 cm | medium 1 to 2 cm | medium 2 to 5 mm | medium 2 to 5 mm |
| | thick 5 to 10 mm | coarse 5 to 10 cm | coarse 5 to 10 cm | coarse 2 to 5 cm | coarse 2 to 5 cm | coarse 5 to 10 mm | not found |
| | very thick > 10 mm | very coarse > 10 cm | very coarse > 10 cm | very coarse > 5 cm | very coarse > 5 cm | very coarse > 10 mm | not found |

Fig. 1.6 Characteristics of various kinds of soil peds.

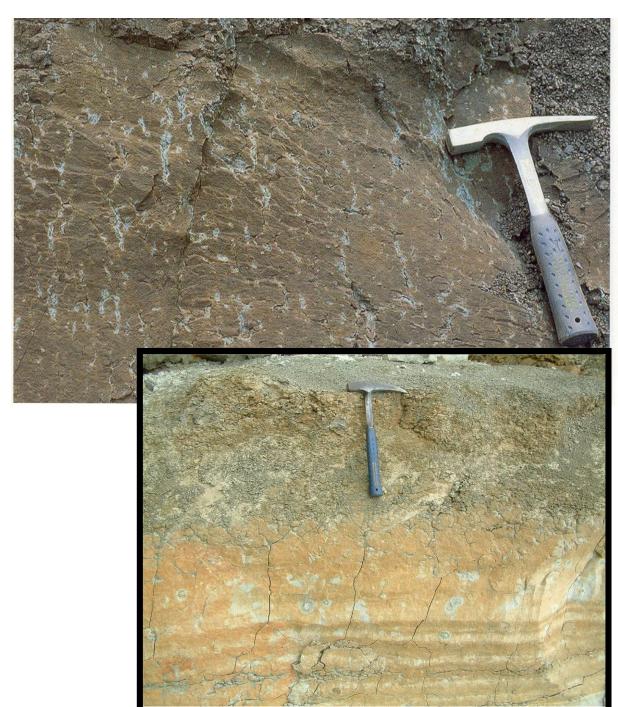
Colour photo 38. **Blocky peds.** The slicken-sided purplish black surfaces of this rock specimen are former cutans of iron-manganese oxides (mangans) which define a coarse angular blocky ped. The greenish grey interior of the ped is visible to the upper right, where it has broken open. This rock specimen is from the gleyed subsurface (Bg) horizon of an Aquept paleosol in the late Eocene (Duchesnean North American Land Mammal "Age"), upper Clarno Formation in the red hill west of Hancock Field Station, near Clarno, Oregon, USA (Retallack, 1991c,d). Scale is in centimetres and millimetres.

Colour photo 39. Slickensided clay skins. The randomly arranged striated surfaces of this rock specimen are compacted clay skins which define very coarse angular blocky peds. The grey/pink mottling is probably due to burial gleization and hydroxide dehydration during burial of a formerly yellow-brown soil. This rock specimen is from the upper Bt horizon of the Interior clay paleosol, a Paleudalf from the late Eocene (Chadronian North American Land Mammal "Age"), Chadron Formation, in the Pinnacles area of Badlands National Park, South Dakota, USA (Retallack, 1983a). Scale is in millimetres.





Colour photo 40. **Prismatic peds.** The light coloured layer around the hammer is vertically jointed from abundant prismatic peds. This paleosol and others like it supported mangal vegetation within a large tide-dominated delta. Brackish water molluscs have been found in them (Colour photo 21). The prismatic peds are in a Bw horizon beneath a dark carbonaceous A horizon in a Sulfaquent paleosol from the mid-Cretaceous (Cenomanian), upper Dakota Formation, along the Saline River north of Russell, Kansas, USA (Retallack & Dilcher, 1981a,b). The hammer handle for scale is 25 cm long.



Retallack (1997)

Colour photo 5. **Drab-haloed root traces.** Drab-haloes of light green silt surround deeply penetrating root traces and are conspicuous in the red oxidized matrix of the paleosol. They probably formed early during burial of the soil as the root decomposed anaerobically. The paleosol is an Eutrochrept in the Late Devonian (Frasnian), Oneonta Formation, Catskill magnafacies, in road cuts along interstate highway I88 near Unadilla, New York, USA (Gordon & Bridge, 1987). Hammer for scale has a handle 25 cm long.

Pedogenic carbonates (calcrete) from the Upper Carnian Raibl Formation, Dolomites, Southern Alps (Italy) showing columnar internal structures in a green to reddish mudstone matrix reflecting a change in paleogroundwater table.

