SOURCE PULSE ENHANCEMENT BY DECONVOLUTION OF AN EMPIRICAL GREEN'S FUNCTION

Charles S. Mueller*

U.S. Geological Survey, Office of Earthquakes, Volcanoes, and Engineering, Menlo Park, CA

source-time function are enhanced if path, record-
ing-site, and instrument complexities can be re**ing-site, and instrument complexities can be re- is only necessary that the Green's function corearthquake has a simple source, its seismogram can [e.g. Kanamori, 1979]. The above studies, as well** be treated as an empirical Green's function and deconvolved from the seismogram of a larger and/or **deconvolved from the seismogram of a larger and/or Green's functions [e.g. Heaton, 1982], all share more complex earthquake by spectral division. the need to assume a space-time source behavior** When the deconvolution is well posed, the quotient for the earthquake to be modeled. In contrast **spectrum represents the apparent source-time func- the apparent source-time function is obtained tion of the larger event. This study shows that directly using the method outlined in this paper. with high-quality locally recorded earthquake data It is conceptually straightforward to isolate it is feasible to Fourier transform the quotient the source-time spectrum of a large/complex event and obtain a useful result in the time domain. In (considered as a point source) by dividing its** practice, the deconvolution can be stabilized by one of several simple techniques.

tested on high-quality digital recordings of tral California earthquakes. As the present study aftershocks of the Jan. 9, 1982 Miramichi (New shows, with high-quality data the quotient spec-Brunswick) earthquake. In particular, seismograms trum can be transformed to the time domain, yield-
from a Jan. 17 aftershock (017 13:33 GMT, local ing the apparent far-field source-time function of **from a Jan. 17 aftershock (017 13:33 GMT, local ing the apparent far-field source-time function of** mag.=3.5) exhibit path or site effects which com-
plicate the determination of source parameters. **plicate the determination of source parameters. have advantages over amplitude spectra alone in of this event is a simple pulse in displacement can be attached to the result by assuming linear-**

pirical Green's functions by Hartzell [1978] in less than the corner frequency of the Green's synthesizing strong-motion records from a complex function (where the delta-function assumption earthquake. For the purposes of Hartzell's study, holds) the result represents the source-time func-
several subsequent studies *[e.g. Kanamori, 1979;* tion of the large/complex earthquake in units of **several subsequent studies [e.g. Kanamori, 1979; tion of the large/complex earthquake in units of Hadley and Helmberger, 1980; Irikura, 1983; Fran- the seismic moment of the Green's function event. kel and Kanamori, 1983] and this paper, a useful Equivalently, the result represents a scaled verempirical Green's function is a seismogram written sion of the displacement seismogram which the** by a simple earthquake which includes the res-
ponses of path, recording site, and instrument: homogeneous whole-space. ponses of path, recording site, and instrument:

$$
G(t)=S(t) * P(t) * R(t) * I(t), where S(t) \approx \delta(t).
$$

constant or at least that differences are well termined using spectral division to deconvolve an known. Then, assuming that a small earthquake has independently determined source wavelet from telea simple source-time function, the question of its usefulness as an empirical Green's function reduces to how well its seismogram captures the of study, source-time functions of large essential aspects of path and recording site rele- earthquakes were isolated using time-domain vant to the large/complex earthquake under study. deconvolution of simple theoretical Green's func-
Practically speaking, the answer to this question tions [e.g. Boatwright, 1980; Kikuchi and Kana-Practically speaking, the answer to this question involves earthquake locations and focal mechan**involves earthquake locations and focal mechan- mori, 1982; Ruff and Kanamori, 1983]. Here I ap-**

In modeling near-source high-frequency strong ground motion, it may be appropriate to use different Green's functions for different parts of a

Abstract. Observations of the earthquake large fault [e.g. Hartzell, 1978; Irikura, 1983]. respond to an average path and recording site [e.g. Kanamori, 1979]. The above studies, as well

function. Bakun et al. [1976] used this technique
to examine the spectral amplitudes of small cen-In this paper, the method is implemented and to examine the spectral amplitudes of small cen-

ited on high-quality digital recordings of tral California earthquakes. As the present study **Algeright Source modeling. In either domain, meaning**
can be attached to the result by assuming linear**with duration • 0.07 second for both P and S. ity and assuming that the far-field source-time function of the small earthquake is a delta func-Introduction tion in displacement. The spectral division serves to deconvolve path, recording-site, and** Small-earthquake seismograms were used as em- instrument complexities. Then for frequencies
'ical Green's functions by Hartzell [1978] in less than the corner frequency of the Green's

In earthquake seismology, deconvolution techniques have primarily been used in two areas of teleseismic-body-wave research. In one type of study, upper-mantle or crustal structure was de-I assume in what follows that instrumentation is study, upper-mantle or crustal structure was de-
constant or at least that differences are well termined using spectral division to deconvolve an 1971; Clayton and Wiggins, 1976]. In another type
of study, source-time functions of large ply the spectral-division technique to relatively high-frequency local earthquake data; the main objective being the determination of the source function. Divisional deconvolution does not re**quire any assumptions about the empirical Green's 1976] and also provides a useful intermediate** This paper is not subject to U.S. copyright. Pub-

1ished in 1985 by the American Geophysical Union.
the empirical Green's function also makes the **lished in 1985 by the American Geophysical Union. the empirical Green's function also makes the method robust seismologically; if the Green's Paper number 4L6368. function is well chosen, no assumptions need to be**

^{*} also at Dept. of Geophysics, Stanford Univ.