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## Marine hyperpycnal flows: initiation, behavior and related deposits. A review

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### Abstract

Hyperpycnal flows form in the marine environment when river discharge enters the ocean with suspended concentrations in excess of  $36 \text{ kg m}^{-3}$  due to buoyancy considerations, or as little as  $1\text{--}5 \text{ kg m}^{-3}$  when convective instability is considered. They form at a river mouth during floods of small to medium size rivers including extreme events such as jökulhups, dam breaking and draining, and lahars. Associated with high-suspended concentration, they can transport considerable volume of sediment to ocean basins. The typical deposit or hyperpycnite sequence is a compound of a basal coarsening-up unit, deposited during the waxing period of discharge, and a top fining-up unit deposited during the waning period of discharge. Hyperpycnites differ from other turbidites because of their well-developed inversely graded facies and intrasequence erosional contacts. These observations lead to a complete redefinition and interpretation of fine-grained turbidites. Hyperpycnite stacking can locally generate high-sedimentation rates, in the range of 1–2 m per 100 year. Because hyperpycnites are related to climate through flood frequency and magnitude, their record should vary with sea level and climate change. They can also be associated with proximal ice-melting settings. Hyperpycnal flows could also be involved in the formation of meandering canyons and channels.

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### 1. Introduction

In the past 30 years, offshore mass wasting processes and gravity flows have been widely studied, stimulated by the needs of deep-water reservoir characterization and other offshore industries needing protection against natural hazards. The importance of gravity processes is associated with new exploration of deep-marine environments including ODP drilling on the Amazon Fan (Hiscott, Pirmez, & Flood, 1997; Normark & Damuth, 1997) or industrial-academic surveys of the Zaire turbidite system (Savoye et al., 2000).

Following pioneer work of Kuenen and Migliorini (1950), early classifications of offshore gravity processes arose in the 1970s (Lowe, 1982; Middleton, 1976, 1993; Middleton & Hampton, 1973; Nardin, Hein, Gorsline, & Edwards, 1979). These classifications underlined the existence of two major

processes in the marine environment: mass flows and turbidity currents. Mulder and Cochonat (1996) and Shanmugam (1996) discussed the complexity of a classification, since one single event may involve several processes from the failure to final deposition. Turbidity flows (Kneller & Buckee, 2000; Lowe, 1979, 1982; Middleton & Hampton, 1973, 1976; Nardin et al., 1979; Stow, 1996) result from slide transformation, continuation of a fluvial flow or concentration processes. Ignitive transformation of a submarine slide into a flow in which turbulent energy substantively increases (Emms, 1999; Fukushima, Parker, & Pantin, 1985; Parker, 1982; Parker, Fukushima, & Pantin, 1986) has been described worldwide, including the 1929 Grand Banks event (Hughes-Clarke, 1990; Hughes-Clarke, Shor, Piper, & Mayer, 1990; Piper, Cochonat, Ollier Le Drezen, Morrison & Baltzer, 1992), and the 1979 Nice event (Genesseeux, Mauffret, & Pautot, 1980; Malinverno, Ryan, Auffret, & Pautot, 1988; Piper & Savoye, 1993).

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